

Flow in unsaturated porous medium

- **Motion equation**

1. Ignore the movement of water in the vapor phase
2. Ignore the movement due to pressure differences resulting from variations in salt concentration (*osmotic effect*)
3. Ignore the movement due to temperature variations (*thermo-osmotic effect*)
4. assume that the solid matrix is rigid and stable
5. the void space is occupied by a wetting (water) and a non-wetting phase (air)
6. the soil is isotropic

We shall consider flow resulting only from variations in piezometric head or in pressure in the water; the density ρ_w may vary.

The flow of the wetting and non-wetting phases can be described using the Darcy's law modified considering the permeability change owned to the presence of the other phase which occupies part of the void space.

$$q = -K \cdot \nabla \Phi = -\frac{k}{\mu} (\nabla p + \rho g \cdot 1 \cdot z) = -\frac{k}{\mu} \left(\frac{\partial p}{\partial x} \cdot 1 \cdot x + \frac{\partial p}{\partial y} \cdot 1 \cdot y + \left(\frac{\partial p}{\partial z} + \rho g \right) \cdot 1 \cdot z \right)$$

so

wetting phase:
$$q_{iw} = -\frac{k_{ijw}(S_w)}{\mu_w} \left(\frac{\partial p_w}{\partial x_j} + \rho_w g \cdot \frac{\partial z}{\partial x_j} \right) = -k_{ij} \frac{k_{rw}(S_w)}{\mu_w} \left(\frac{\partial p_w}{\partial x_j} + \rho_w g \cdot \frac{\partial z}{\partial x_j} \right)$$

non-wetting phase:
$$q_{inw} = -\frac{k_{ijnw}(S_{nw})}{\mu_{nw}} \left(\frac{\partial p_{nw}}{\partial x_j} + \rho_{nw} g \cdot \frac{\partial z}{\partial x_j} \right) = -k_{ij} \frac{k_{rnw}(S_{nw})}{\mu_{nw}} \left(\frac{\partial p_{nw}}{\partial x_j} + \rho_{nw} g \cdot \frac{\partial z}{\partial x_j} \right)$$

$$i = 1, 2, 3$$

where:

k_{ijw} : unsaturated permeability of the wetting phase

k_{ijnw} : unsaturated permeability of the non-wetting phase

k_{ij} : permeability at saturation

k_{rw} : relative permeability of the wetting phase

k_{rnw} : relative permeability of the non-wetting phase

- **Mass conservation equation**

wetting phase:
$$\frac{\partial(\rho_w n S_w)}{\partial t} + \nabla \cdot (\rho_w q_w) = 0$$

non-wetting phase:
$$\frac{\partial(\rho_{nw} n S_{nw})}{\partial t} + \nabla \cdot (\rho_{nw} q_{nw}) = 0$$

- **Continuity equation**

The subsequent equations, valid for the

wetting phase:

$$\frac{\partial(\rho_w n S_w)}{\partial t} - \frac{\partial}{\partial x_i} \cdot \left[\rho_w k_{ij} \frac{k_{rw}(S_w)}{\mu_w} \left(\frac{\partial p_w}{\partial x_j} + \rho_w g \cdot \frac{\partial z}{\partial x_j} \right) \right] = 0$$

non-wetting phase:

$$\frac{\partial(\rho_{nw} n S_{nw})}{\partial t} - \frac{\partial}{\partial x_i} \cdot \left[\rho_{nw} k_{ij} \frac{k_{rnw}(S_{nw})}{\mu_{nw}} \left(\frac{\partial p_{nw}}{\partial x_j} + \rho_{nw} g \cdot \frac{\partial z}{\partial x_j} \right) \right] = 0$$

respectively, define the problem of the flow into an unsaturated porous medium.

The model is completed by a condition on the saturations of the wetting and non-wetting phases:

$$S_w + S_{nw} = 1$$

the relationships between the relative permeabilities of the wetting phase and the non-wetting one with the respective saturations

$$\begin{aligned} k_{rw} &= k_{rw}(S_w) \\ k_{rnw} &= k_{rnw}(S_{nw}) \end{aligned}$$

and of the moisture retention curve $S_w = S_w(p_c)$ where $p_c = p_{nw} - p_w$

Now let us simplify the problem considering only the motion in a vertical plane x-z and suppose that:

1. The non-wetting phase is at $p_{nw} = 0$
2. $\rho_w = \text{const}$

The non-wetting phase doesn't move anymore and the motion equation for the wetting phase becomes:

$$q_w = -K_{Sw} k_{rw}(S_w) \nabla(\Psi_w + z)$$

where $\Psi_w = \frac{p_w}{\rho_w g}$

that is:

direction x:

$$q_{xw} = -K_{Sw} k_{rw}(S_w) \left(\frac{1}{\rho_w g} \frac{\partial p_w}{\partial x} \right)$$

direction z:

$$q_{zw} = -K_{Sw} k_{rw}(S_w) \left(\frac{\partial \Psi}{\partial z} + 1 \right)$$

where:

$$K_{Sw} = \frac{k_w \rho_w g}{\mu_w} : \text{conductivity of wetting phase at saturation}$$

whereas the continuity equation becomes:

$$\frac{\partial(nS_w)}{\partial t} - \nabla \cdot [K_{S_w} k_{rw}(S_w) \nabla(\Psi_w + z)] = 0$$

that is

$$\frac{\partial(nS_w)}{\partial t} = \frac{\partial}{\partial x} \left[K_{S_w} k_{rw}(S_w) \frac{\partial \Psi_w}{\partial x} \right] + \frac{\partial}{\partial z} \left[K_{S_w} k_{rw}(S_w) \left(\frac{\partial \Psi_w}{\partial z} + 1 \right) \right]$$

If you remind that the relationship between moisture content and saturation in water is: $\theta_w = n \cdot S_w$, the continuity equation for the vertical unsaturated porous media flow becomes:

$$\frac{\partial \theta_w}{\partial t} = \frac{\partial}{\partial x} \left[K_{S_w} k_{rw}(\theta_w) \frac{\partial \Psi_w}{\partial x} \right] + \frac{\partial}{\partial z} \left[K_{S_w} k_{rw}(\theta_w) \left(\frac{\partial \Psi_w}{\partial z} + 1 \right) \right]$$

Recalling that saturation depends on capillary pressure, you can write $\frac{\partial \theta_w}{\partial t}$, generally called *general storage term*, as:

$$\frac{\partial \theta_w}{\partial t} = \frac{\partial \theta_w}{\partial \Psi_w} \cdot \frac{\partial \Psi_w}{\partial t} = \frac{\partial(nS_w)}{\partial \Psi_w} \cdot \frac{\partial \Psi_w}{\partial t} = \left[\frac{\partial n}{\partial \Psi_w} \cdot S_w + \frac{\partial S_w}{\partial \Psi_w} \cdot n \right] \cdot \frac{\partial \Psi_w}{\partial t} = \left[S_s \cdot S_w + \frac{\partial S_w}{\partial \Psi_w} \cdot n \right] \cdot \frac{\partial \Psi_w}{\partial t}$$

Now we can write the equation of vertical flow for an unsaturated medium known also as *Richards' equation*:

$$\left[S_s \cdot S_w + \frac{\partial S_w}{\partial \Psi_w} \cdot n \right] \cdot \frac{\partial \Psi_w}{\partial t} = \frac{\partial}{\partial x} \left[K_{S_w} k_{rw}(\theta_w) \frac{\partial \Psi_w}{\partial x} \right] + \frac{\partial}{\partial z} \left[K_{S_w} k_{rw}(\theta_w) \left(\frac{\partial \Psi_w}{\partial z} + 1 \right) \right]$$